

Inferring CCN Properties of Arctic Haze Layers During 2008 ARCPAC Field Campaign

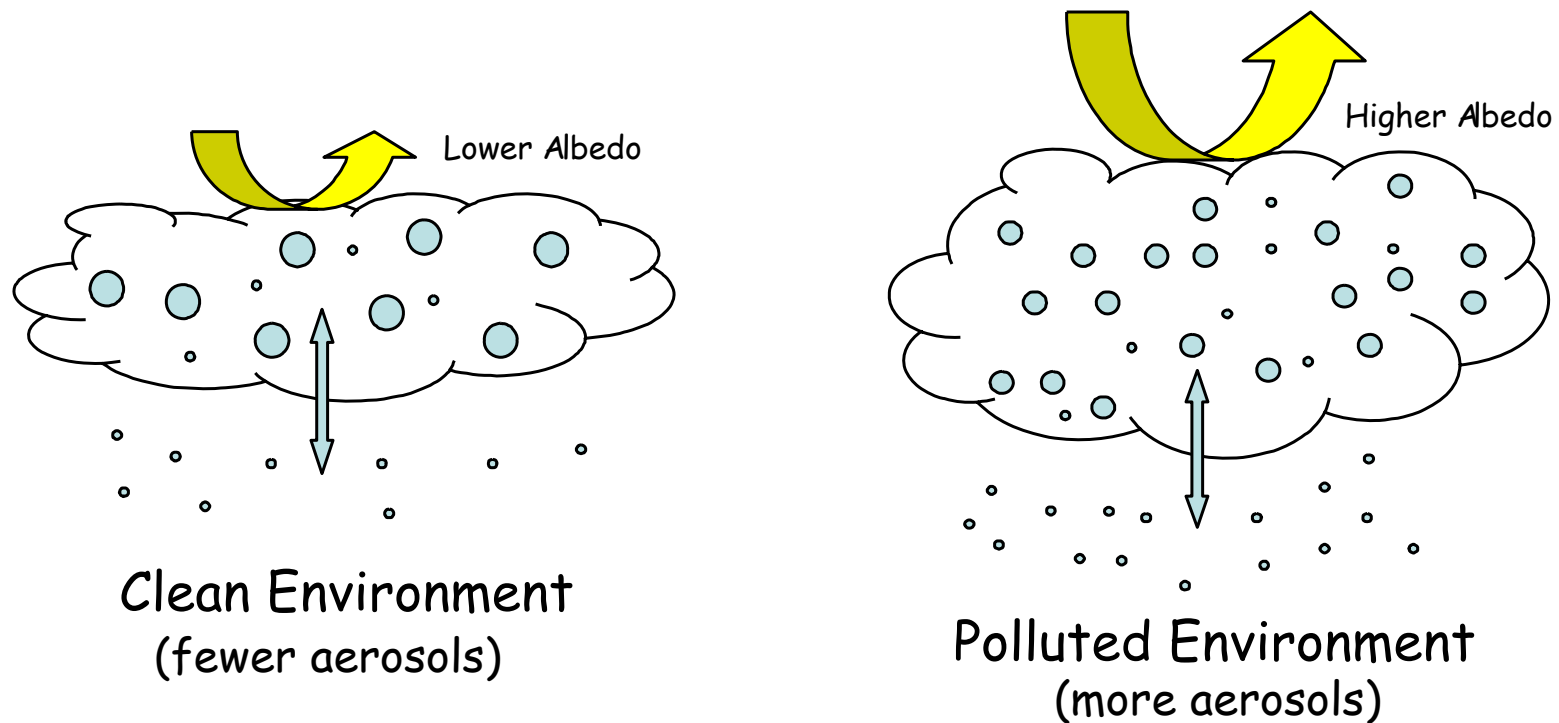
Richard H. Moore
GREF Fellow



Overview

- Motivation: Aerosols, Clouds, Climate
- Cloud Droplet Formation and Growth
- CCN Measurement & Characterization
 - CCN Closure
 - Organic Properties
 - Surface Tension Effects
 - Mixing State
 - Droplet Growth Kinetics
- 2008 ARCPAC Field Campaign
- Vertical Profiles of Haze

Motivation: Aerosol Indirect Effects



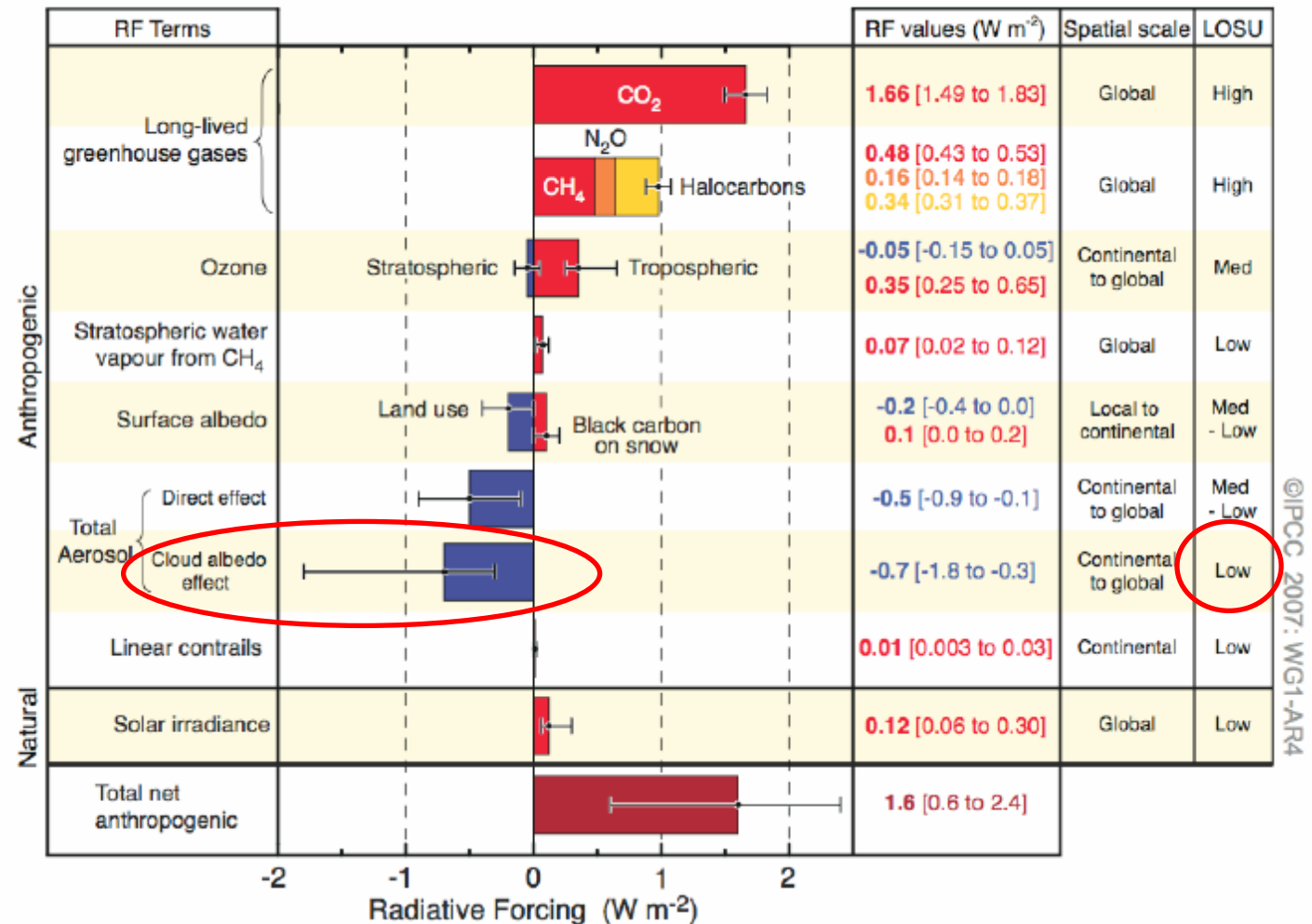
1st Indirect Effect: Polluted clouds will have smaller and more numerous droplets, which makes the cloud denser and able to reflect more incoming radiation

2nd Indirect Effect: Smaller droplets increase cloud lifetime by suppressing precipitation

Motivation: Aerosol Indirect Effects

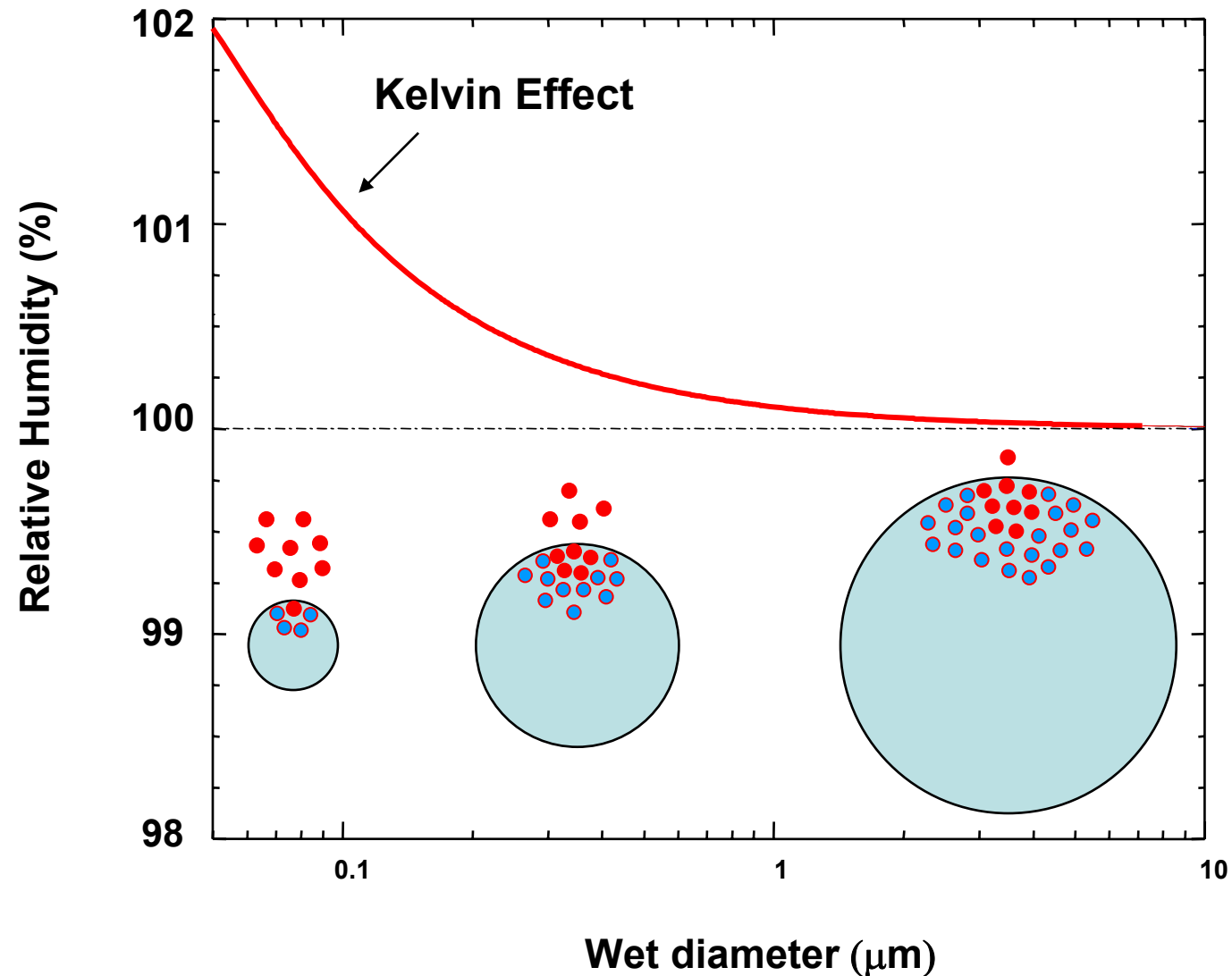
Radiative Forcing Components

- Can compare present-day energy balance to preindustrial (1750 AD).
- Impacts of greenhouse gases are well understood
- Impacts of aerosols are poorly understood



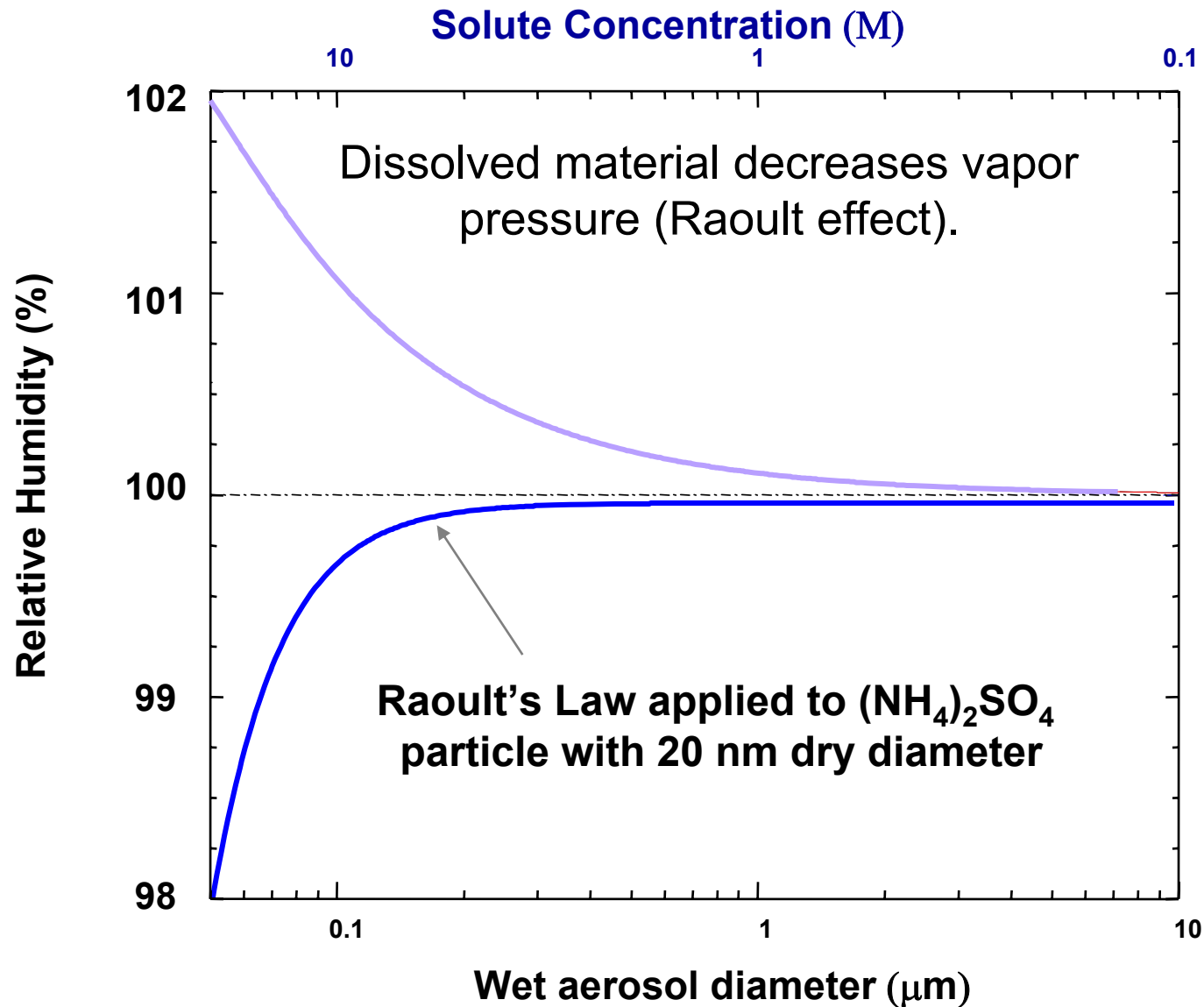
IPCC, 2007

How Do Aerosols Form Droplets?



Credit: A. Nenes

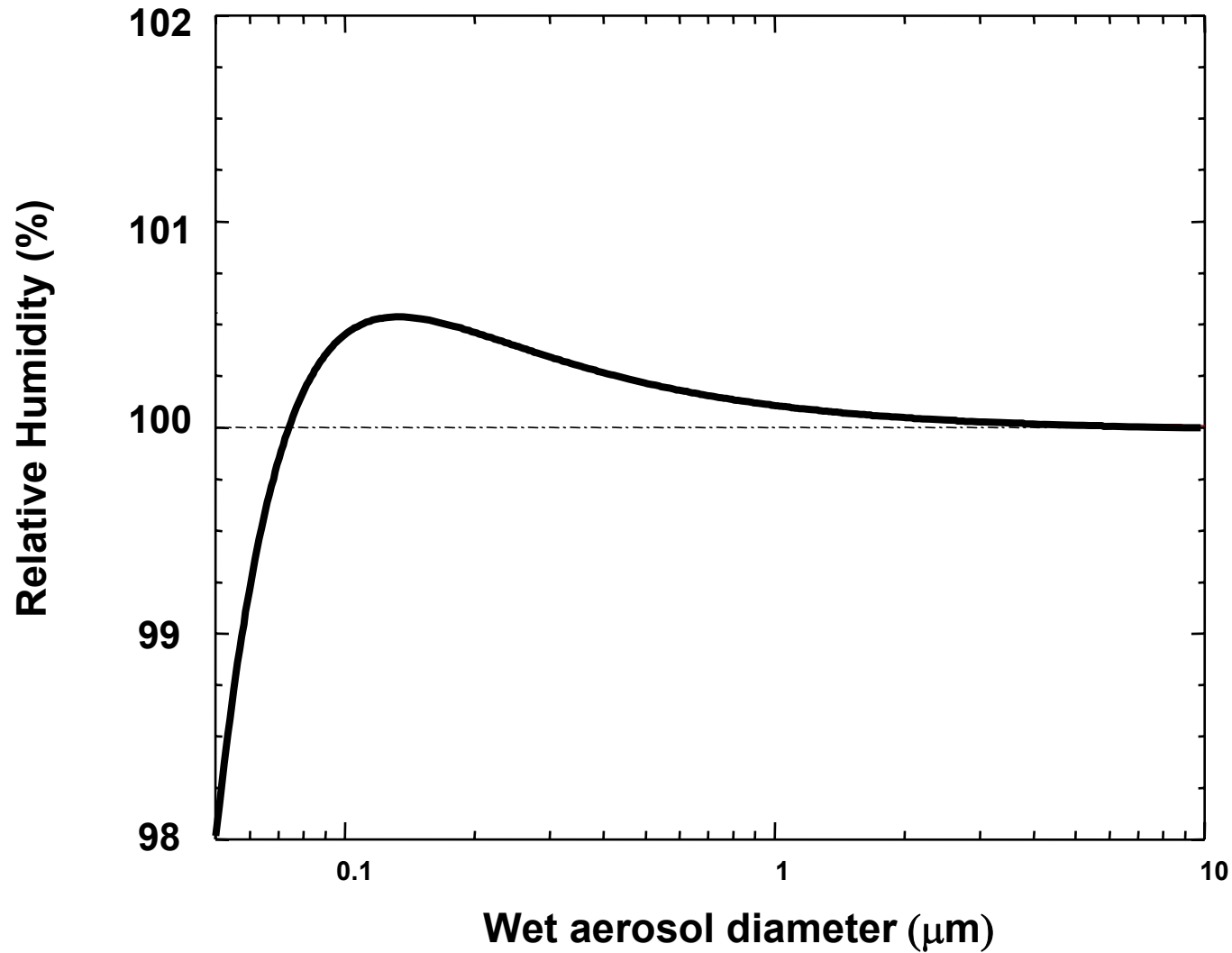
How Do Aerosols Form Droplets?



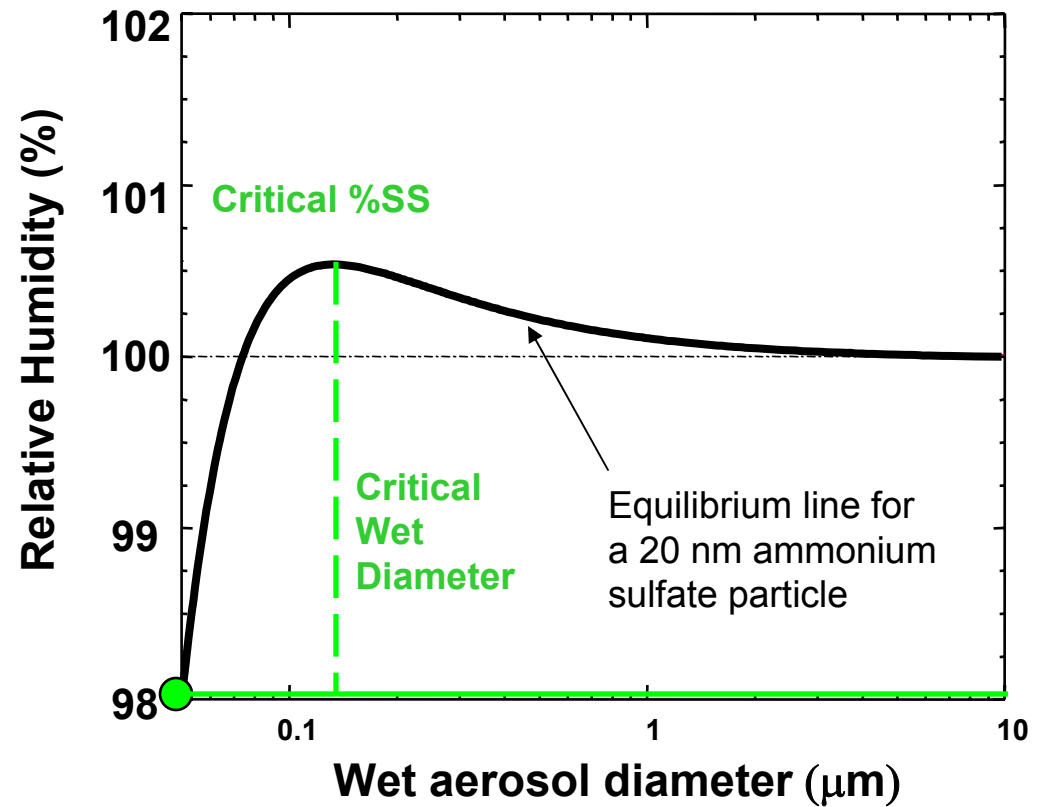
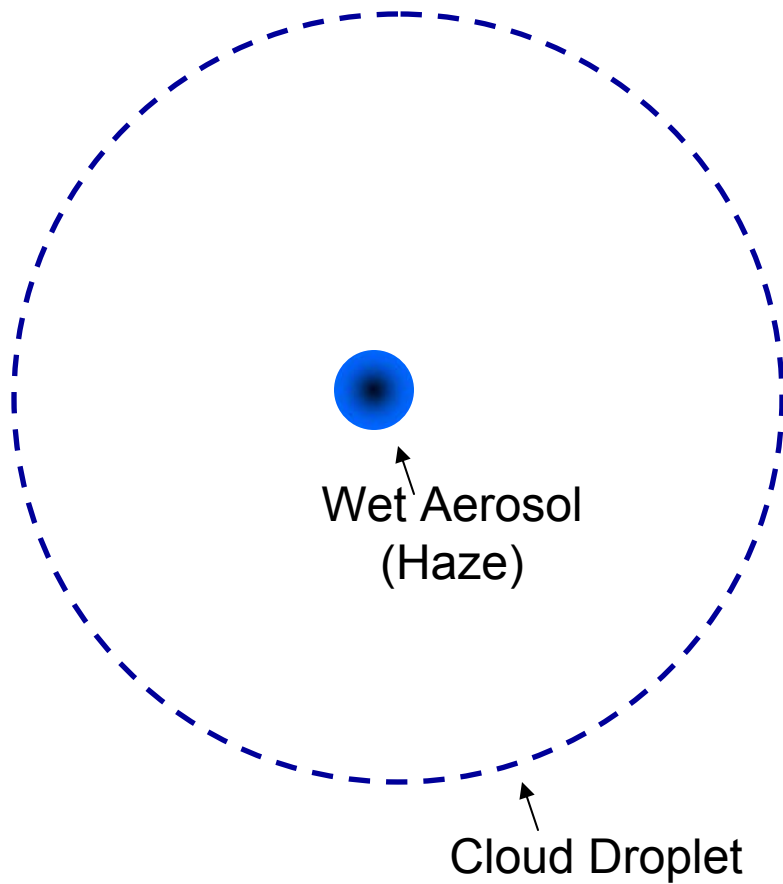
The graph illustrates the relationship between relative humidity and wet aerosol diameter for different theoretical models and a specific chemical species. The y-axis represents Relative Humidity (%) from 98 to 102, and the x-axis represents Wet aerosol diameter (μm) on a logarithmic scale from 0.01 to 10. A secondary x-axis at the top shows Solute Concentration (M) from 10 to 0.1. The Kelvin curve (purple) shows a monotonic decrease in RH with increasing diameter. The Kelvin + Raoult (Köhler Theory) curve (black) shows a characteristic Köhler peak, reaching a maximum RH of approximately 100.5% at a diameter of about 0.15 μm. The Raoult curve (purple) for (NH₄)₂SO₄ with a 20 nm dry diameter shows a rapid increase in RH, approaching 100% at larger diameters. The Köhler equation is provided as follows:

$$\ln\left(\frac{\%RH}{100}\right) \equiv \ln\left(\frac{p_w}{p_w^\circ}\right) = \frac{4M_w\sigma}{RT\rho_w D_p} - \frac{6n_s M_w}{\pi\rho_w D_p^3}$$


How Do Aerosols Form Droplets?

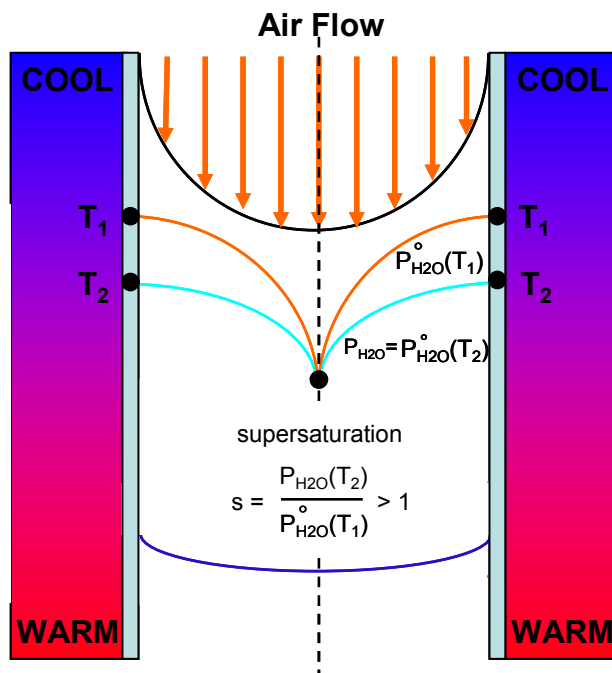


How Do Aerosols Form Droplets?

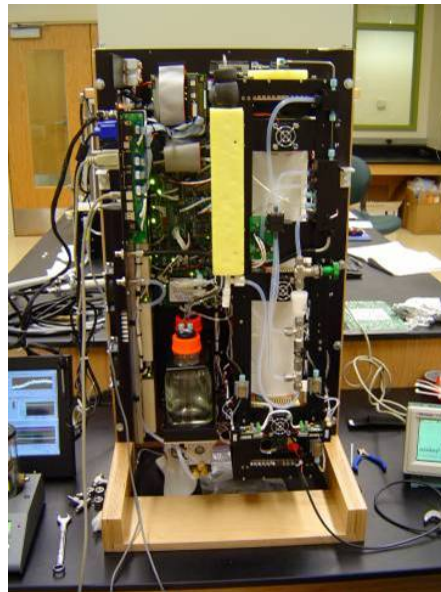


CCN = f (Size, Chemical Composition)

Measuring CCN Activation and Growth with the DMT Continuous-Flow Streamwise Thermal-Gradient CCN Counter



- Expose aerosol to a fixed water vapor supersaturation for a fixed amount of time
- Measure **CCN Concentration** and **Final Droplet Size**.



CCN Closure: Comparing Observations with Predictions from Köhler Theory

Measured **composition** from AMS & Assumptions: Mixing State, Surface Tension, Organic Properties

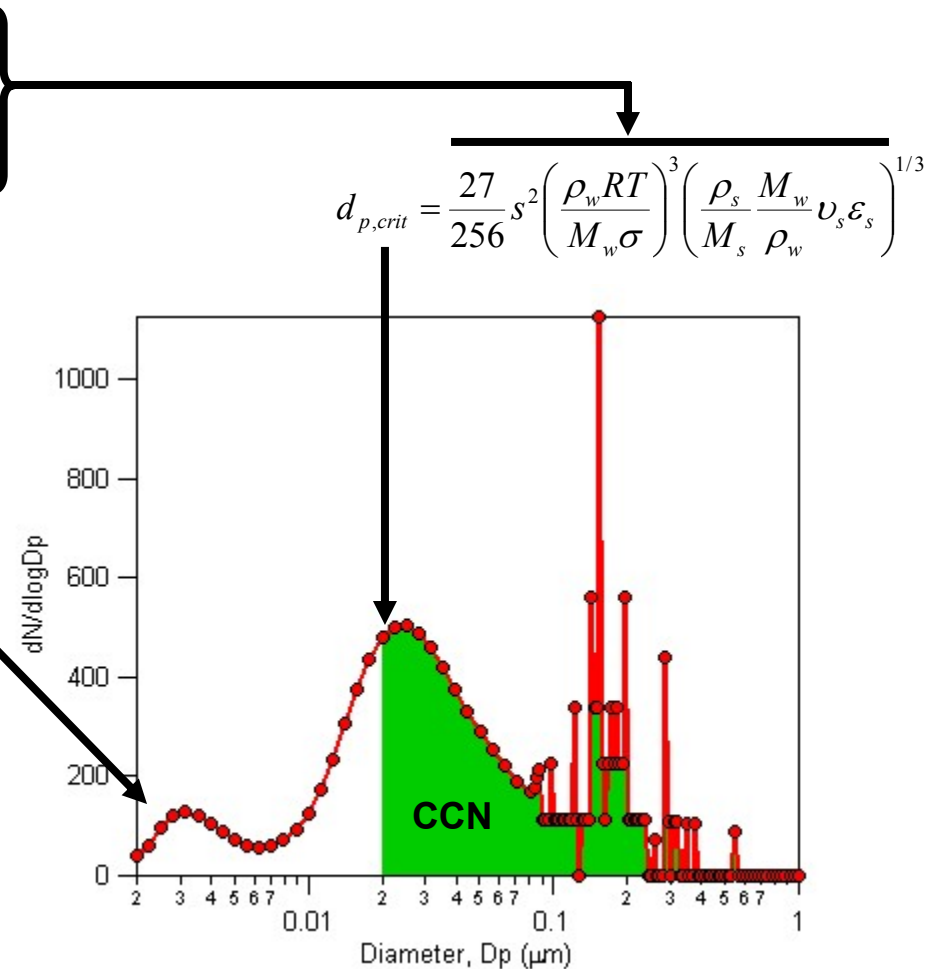
+

Measured **size distribution** from UHSAS, CPCs

=

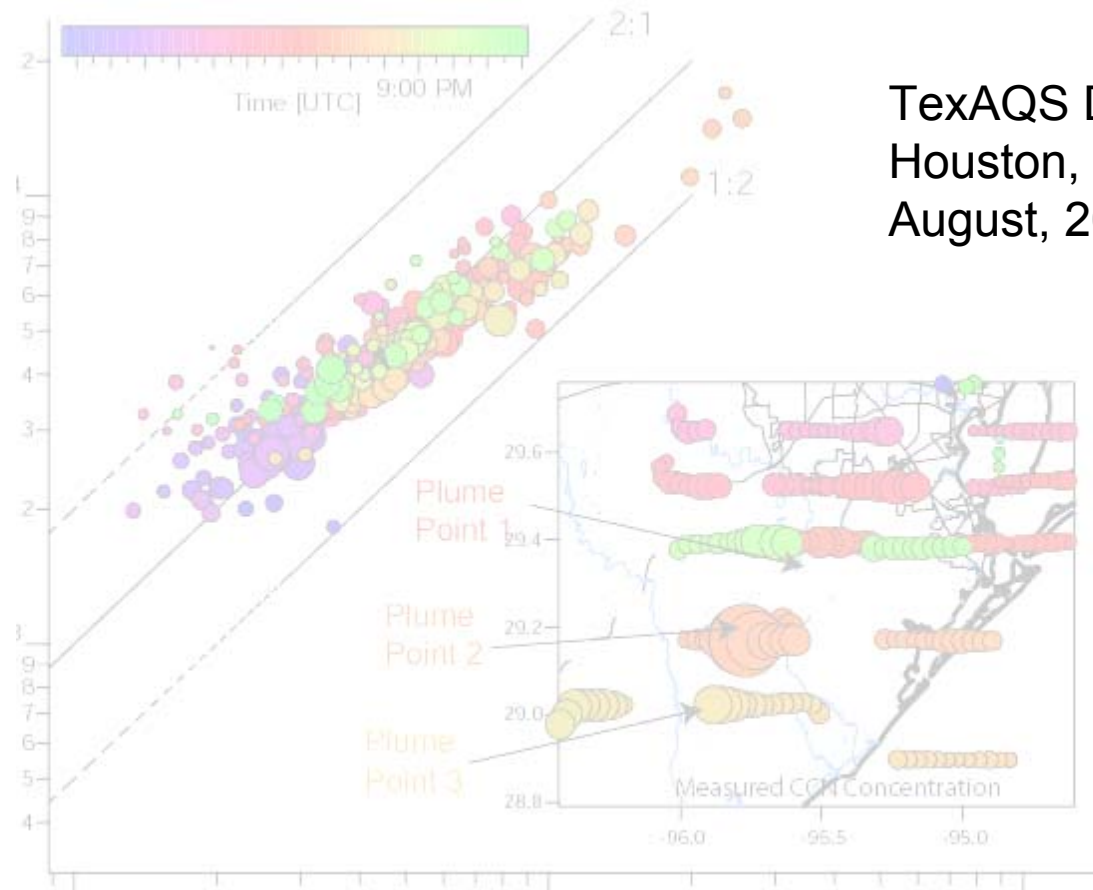
Predicted CCN Concentration

Does the predicted concentration match what we measure?



Sample Size Distribution from TexAQS

CCN Closure: Comparing Observations with Predictions from Köhler Theory

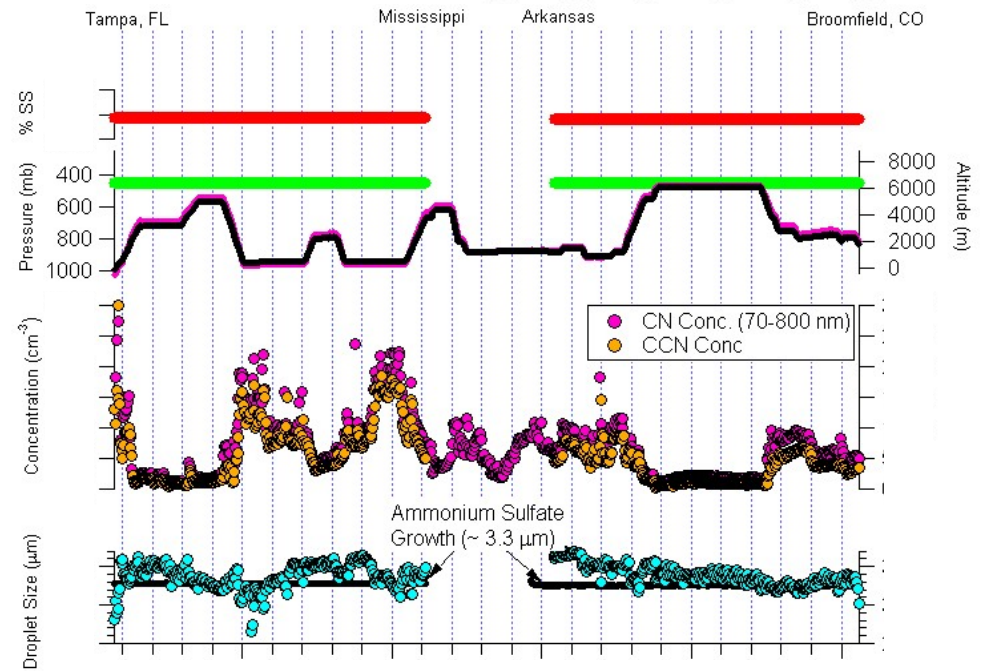
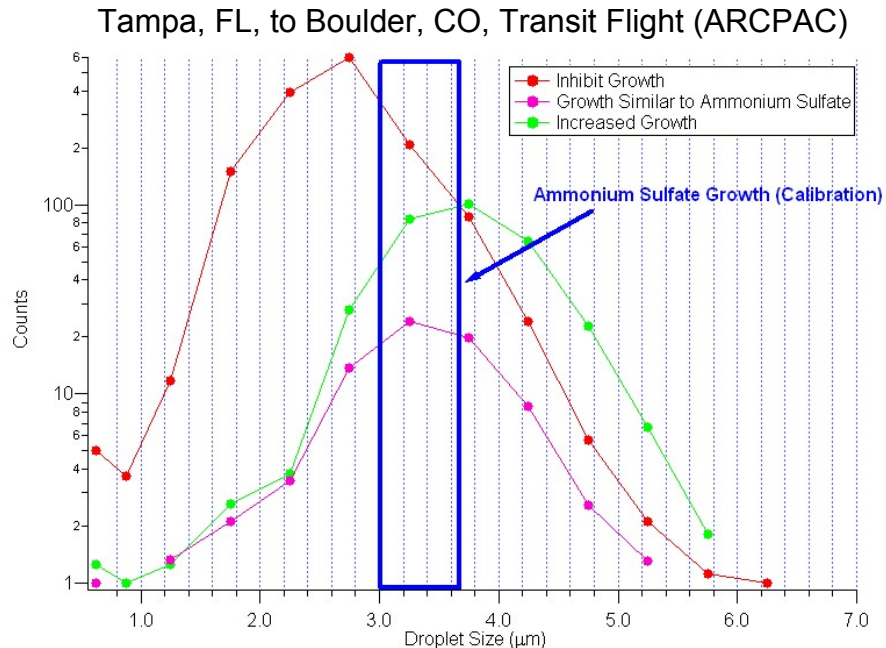
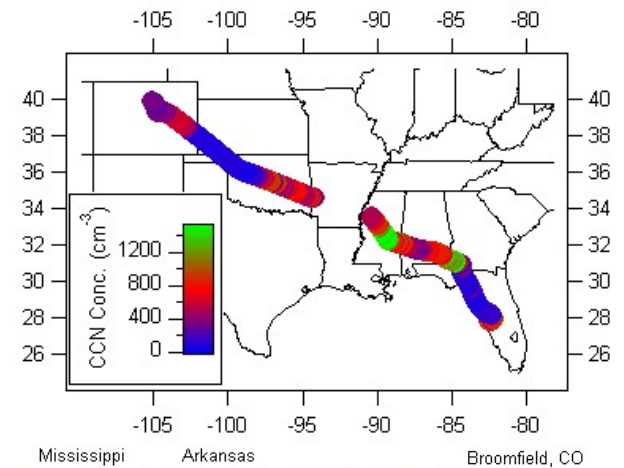


TexAQS Data
Houston, TX
August, 2006

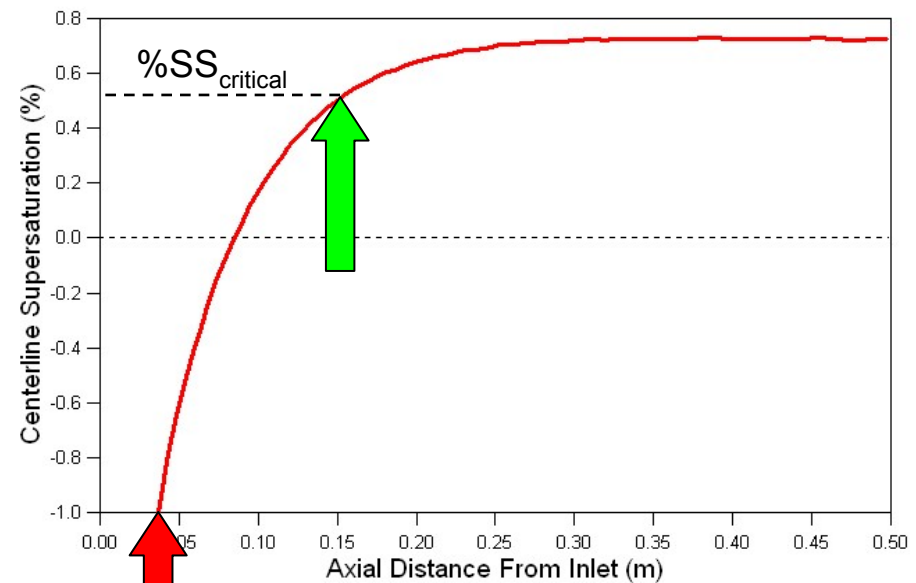
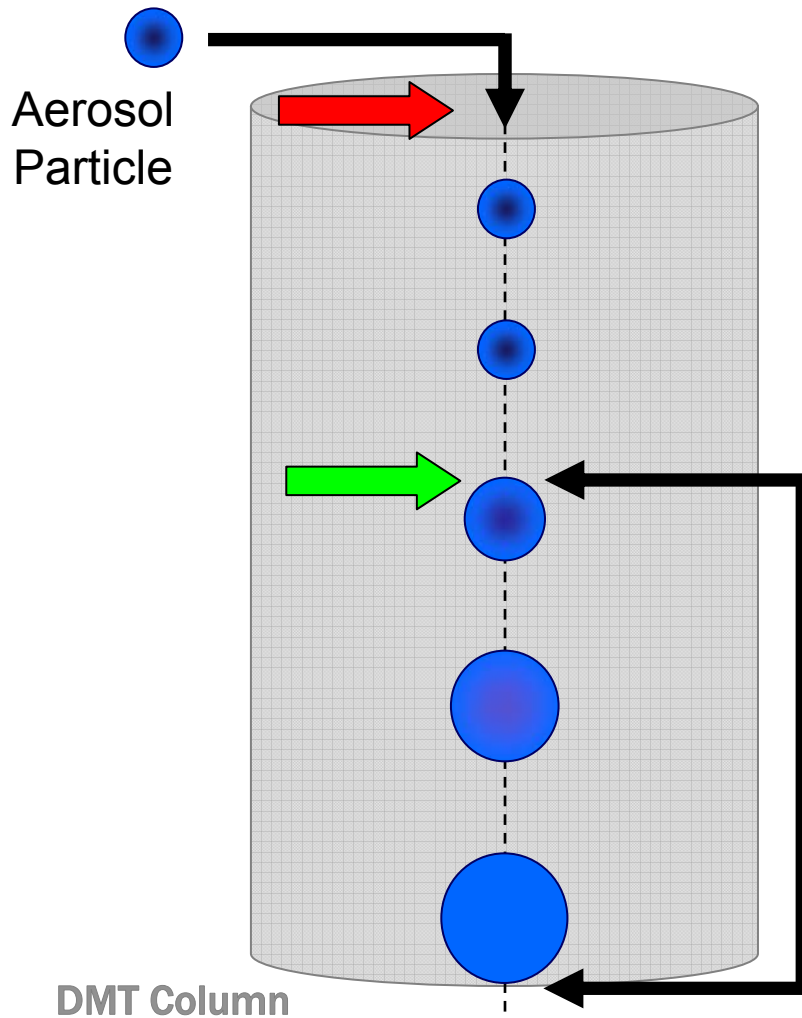
Asa-Awuku et al., *in prep.*

Characterizing Droplet Growth Kinetics

- As a first step, evaluate droplet growth with respect to ammonium sulfate calibration aerosol
- The next step will be to model the aerosol residence time and apply the droplet growth equation to infer the water uptake coefficient



Characterizing Droplet Growth Kinetics



$$\frac{\Delta D_p}{\Delta t} \cong \frac{D_p}{D_p} \cdot \frac{(S_\infty - S_{eq})}{\frac{\rho_w RT_\infty}{4p_w^\circ(T_\infty)D_v' M_w} + \frac{\Delta H_v \rho_w}{4k_a' T_\infty} \left(\frac{\Delta H_v M_w}{T_\infty R} - 1 \right)}$$

$$D_v' = \frac{D_v}{1 + \frac{2D_v}{\gamma_c D_p} \left(\frac{2\pi M_w}{RT} \right)^{1/2}} \quad k_a' = \frac{k_a}{1 + \frac{2k_a}{\gamma_T D_p \rho \hat{c}_p} \left(\frac{2\pi M_a}{RT_a} \right)^{1/2}}$$

Summary of ARCPAC Data

○ FL to CO Transit (3/29) – **POWER FAILURE BETWEEN 17:46 AND 18:38 UTC**

⊘ CO Flight (4/1) – **LEAK**

⊘ CO to AK Transit (4/3) – **INSTRUMENT IN DRYING MODE**

○ AK Flight #1 (4/11) – **POWER FAILURE AFTER 20:46 UTC**

○ AK Flight #2 (4/12)

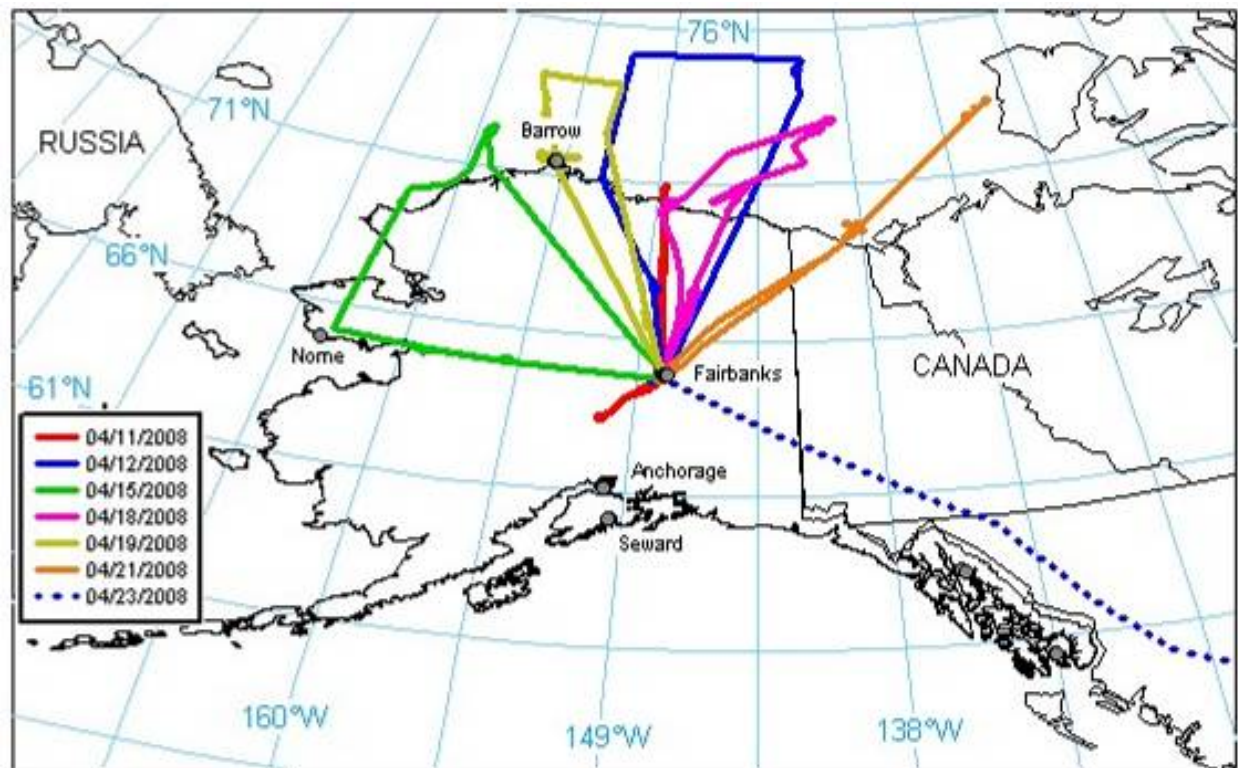
○ AK Flight #3 (4/15)

○ AK Flight #4 (4/18)

○ AK Flight #5 (4/19)

○ AK Flight #6 (4/21)

○ AK to CO Transit (4/23)



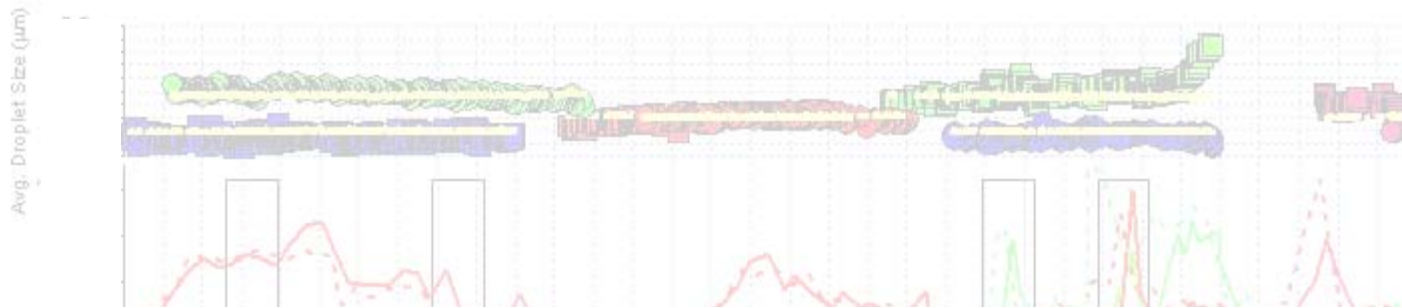
Arctic Haze Layers



from NASA DC-8 (4/12)

Arctic Haze Layers

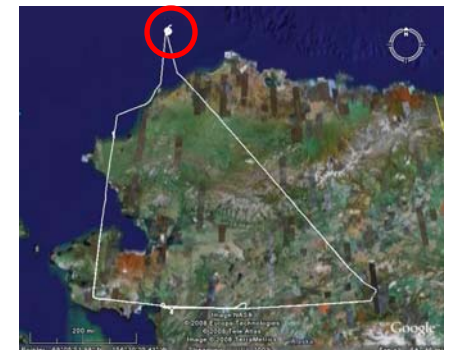
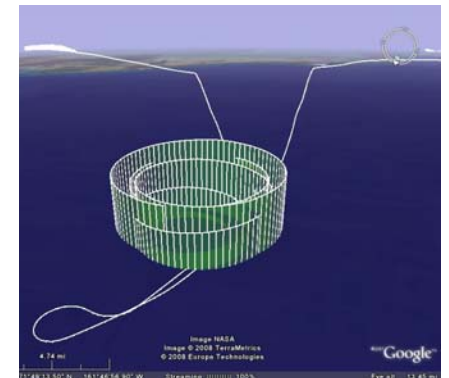
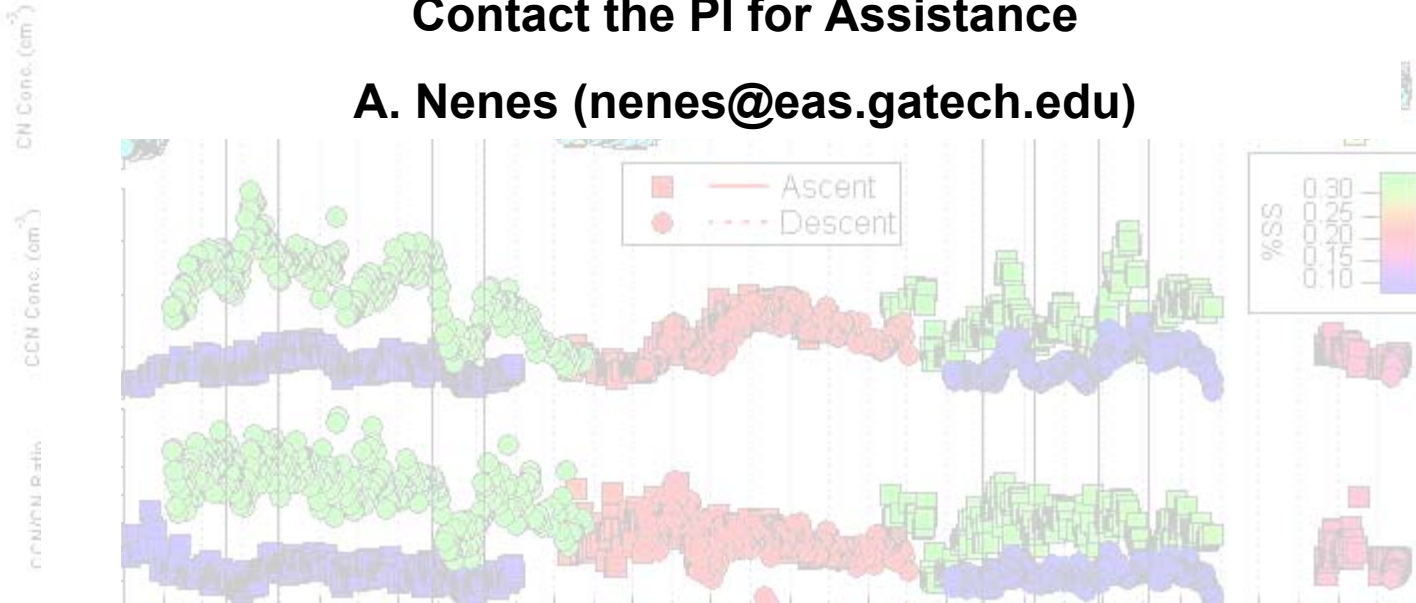
4/15/2008: 2nd Cone top profile: 00:23-00:40 Ascent, 00:40-00:59 Descent



Preliminary CCN Data Omitted Until Finalized

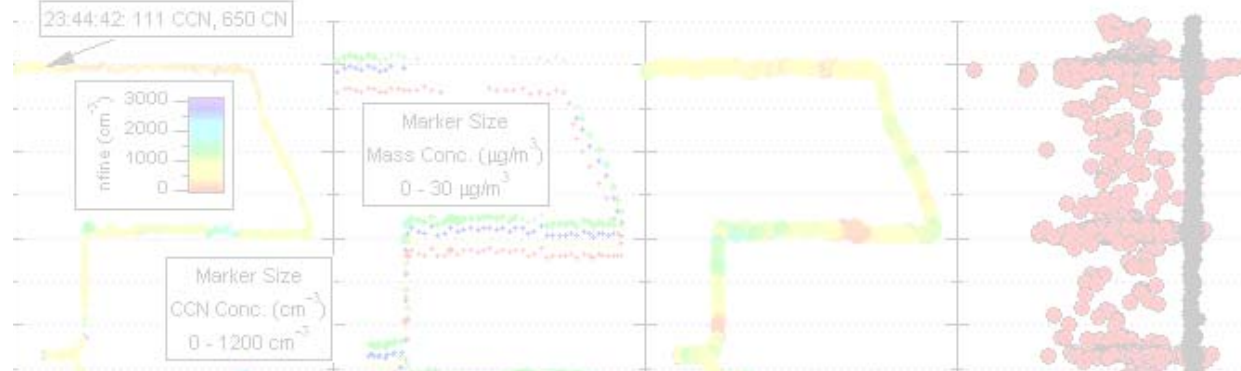
Contact the PI for Assistance

A. Nenes (nenes@eas.gatech.edu)



Arctic Haze Layers

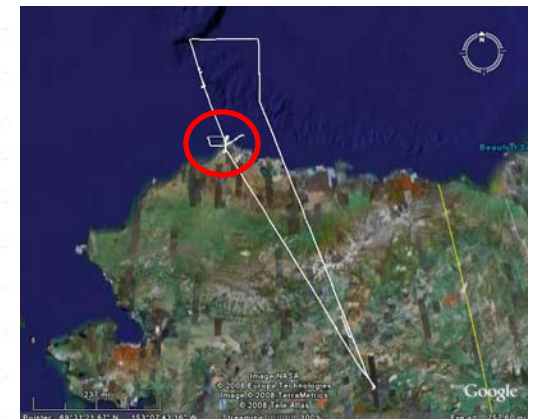
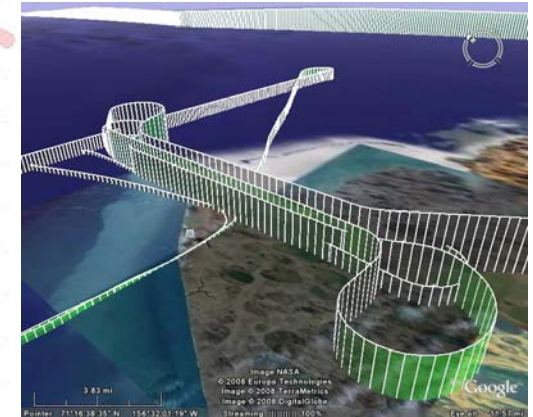
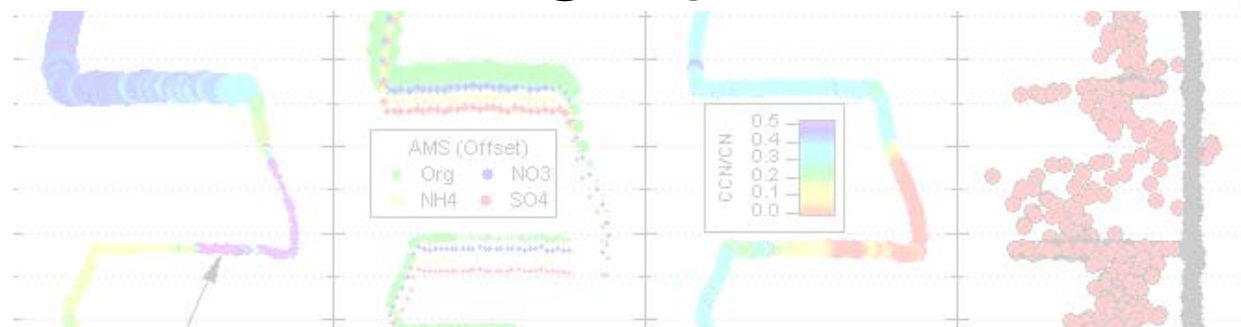
4/19/2008: Stepped profile Over Barrow: 23:30-00:30 Descent



Preliminary CCN Data Omitted Until Finalized

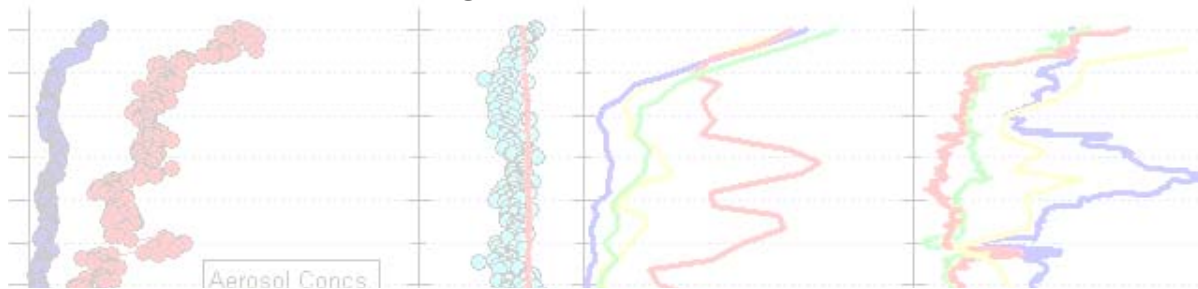
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Arctic Haze Layers

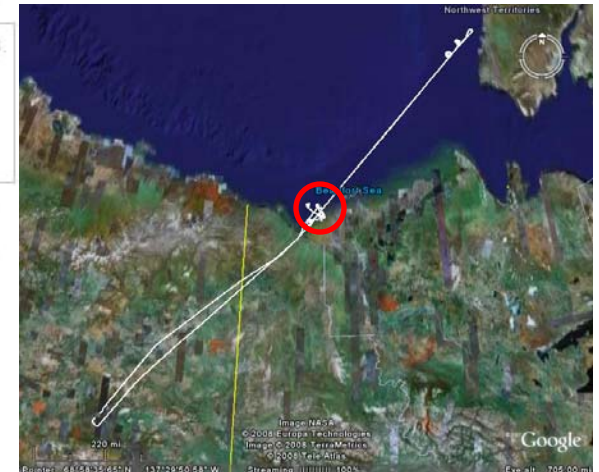
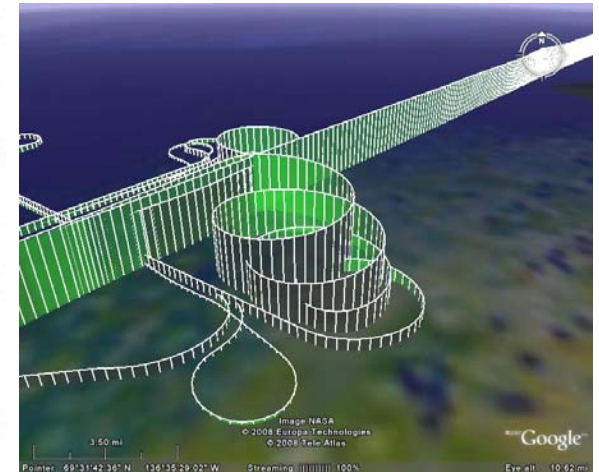
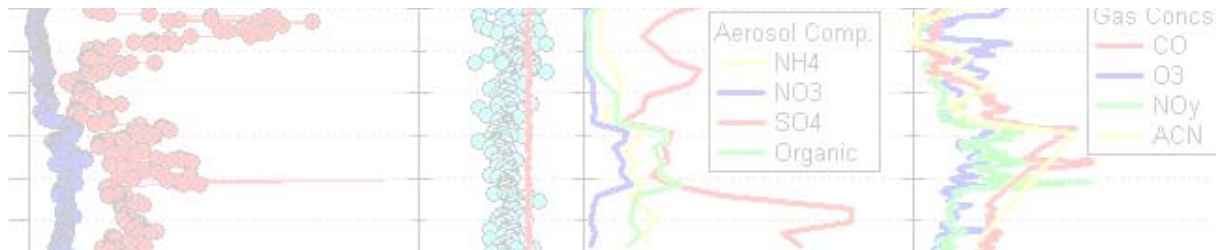
4/21/2008: Descending Profile Near Coast: 23:57-00:08



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Scientific Questions

- Compare CCN Closure & Growth Kinetics:
 - Haze Layer vs. Background
- Do organics contribute to soluble mass?
(Models currently neglect organic contribution to CCN.)
- Do organics depress surface tension and/or lower growth kinetics?
- How can we link inferred CCN properties to observed cloud properties?

Acknowledgements

- Thanos Nenes and the GA Tech Aerosol-Cloud Research Group
- Greg Huey, Dave Tanner, and Mark Lord for Integration Assistance
- NOAA Folks: Chuck Brock, Tom Ryerson, Jeff Peischl, Sara Lance, Julie Cozic, Roya Bahreini, Ann Middlebrook, Joost deGouw, Carsten Warneke, Ken Aiken
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Questions?

